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Introduction

In the sedimentary record of the Transylvanian Paleogene Basin (TPB), Romania, two greater continental-marine sedimentary cycles were distinguished from historical times. The first cycle (Jibou Fm.-Calea Group) is characterized by greater thicknesses and coarser sediments, and as a consequence by stronger tectonic activities. The second cycle (Valea Nădesului Fm. - Turea Group) is less affected by tectonic inversions, permits greater influence to the eustatic control of the accommodation (e.g. contemporaneity of the Brebi Marls Condensed Section with a European maximum of sea-level rise, this work).

The existence of Upper Cretaceous-Paleogene piggy-back basins (e.g. Puini basin, Tarnave basin), within the subsurface of the Transylvanian Depression, were suggested by seismic sections published by Ciulav (2000).

The investigated sedimentary succession - the Turea Group deposits - are supposed to be sedimented in a more eastern member of the piggy-back thrust sequence. The spatial distribution of the transgressive deposits is mainly driven by the inherited topography and is autocyclically controlled, while allocyclic can be more visible close to the Condensed Section.

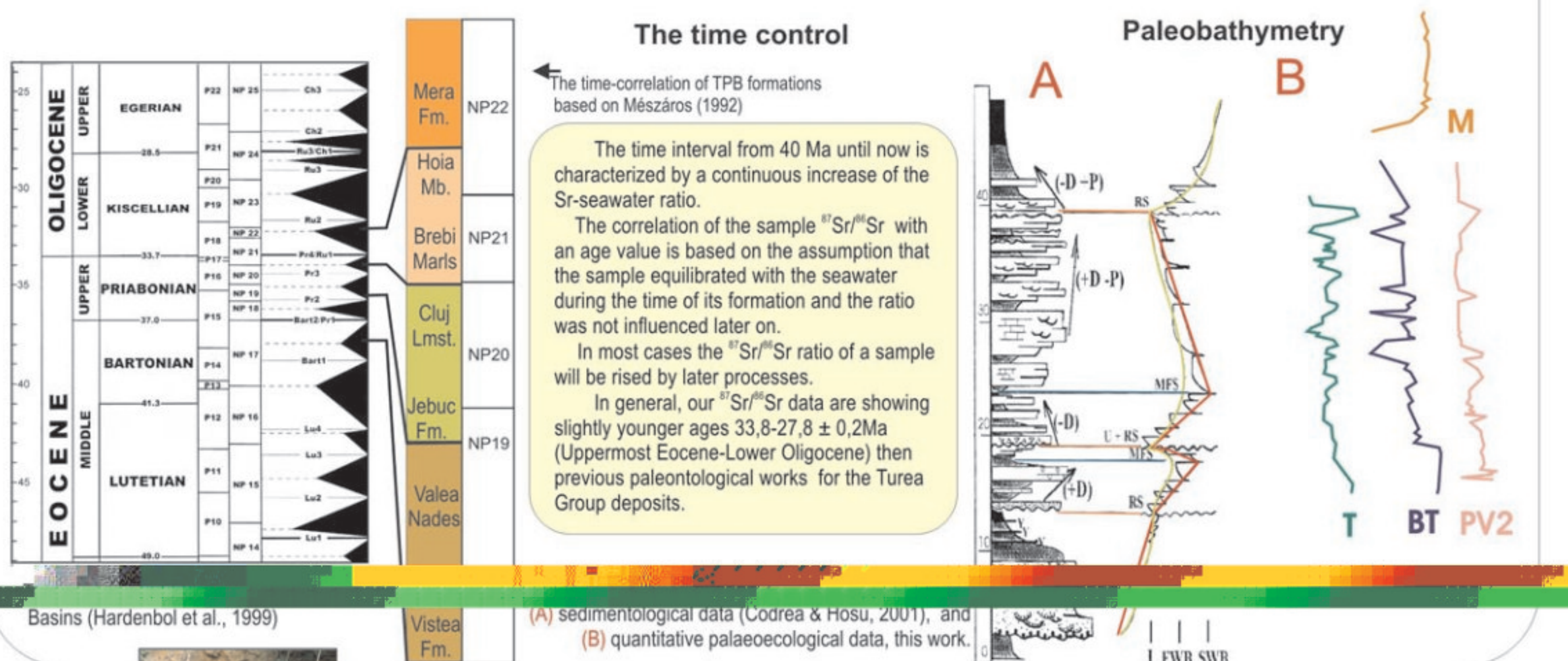


Geological map of the Eocene-Middle Oligocene formations in the Northern Gilau Area (based on Popescu (1984), Raileanu (1956))

Sites of investigation

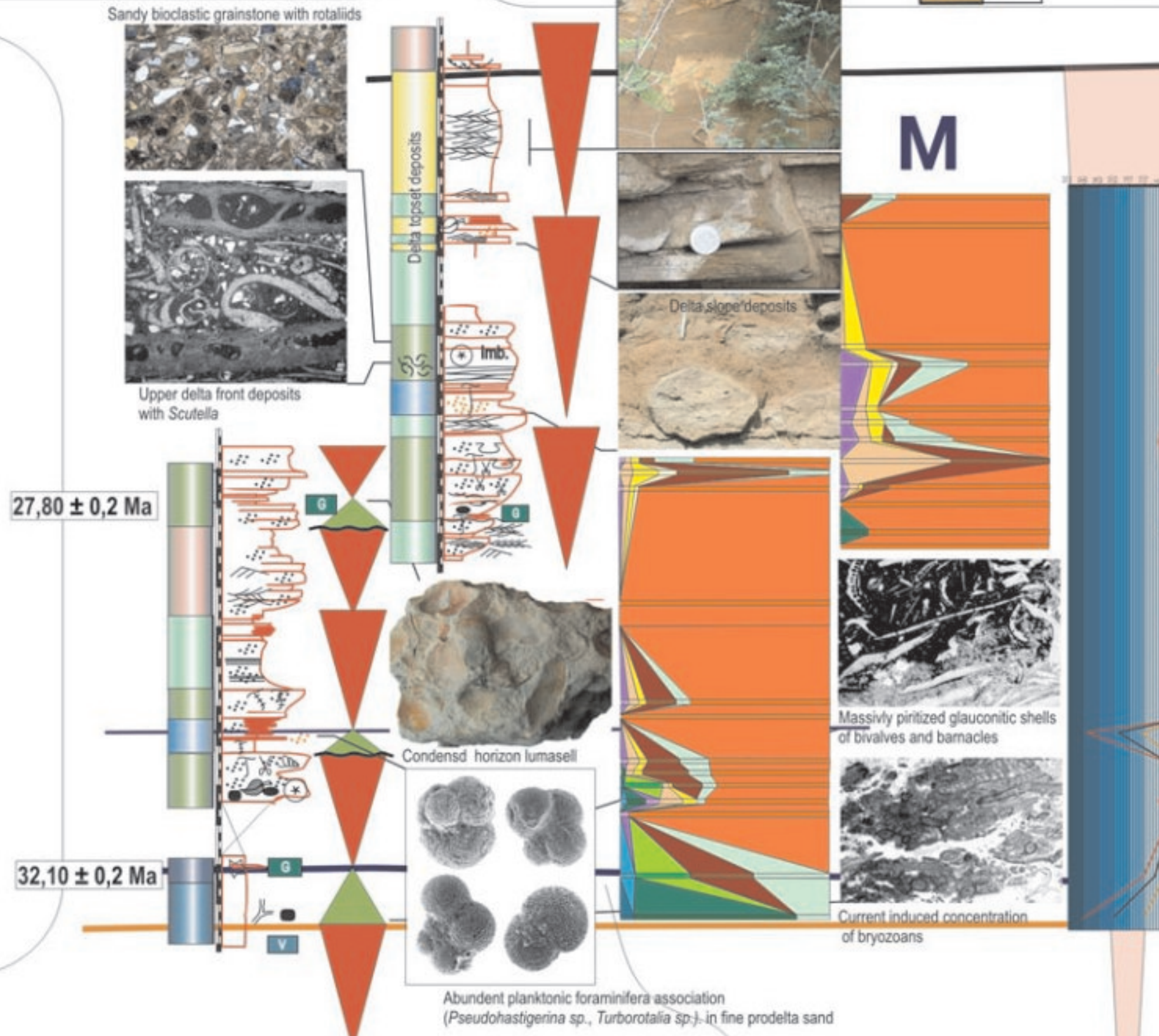
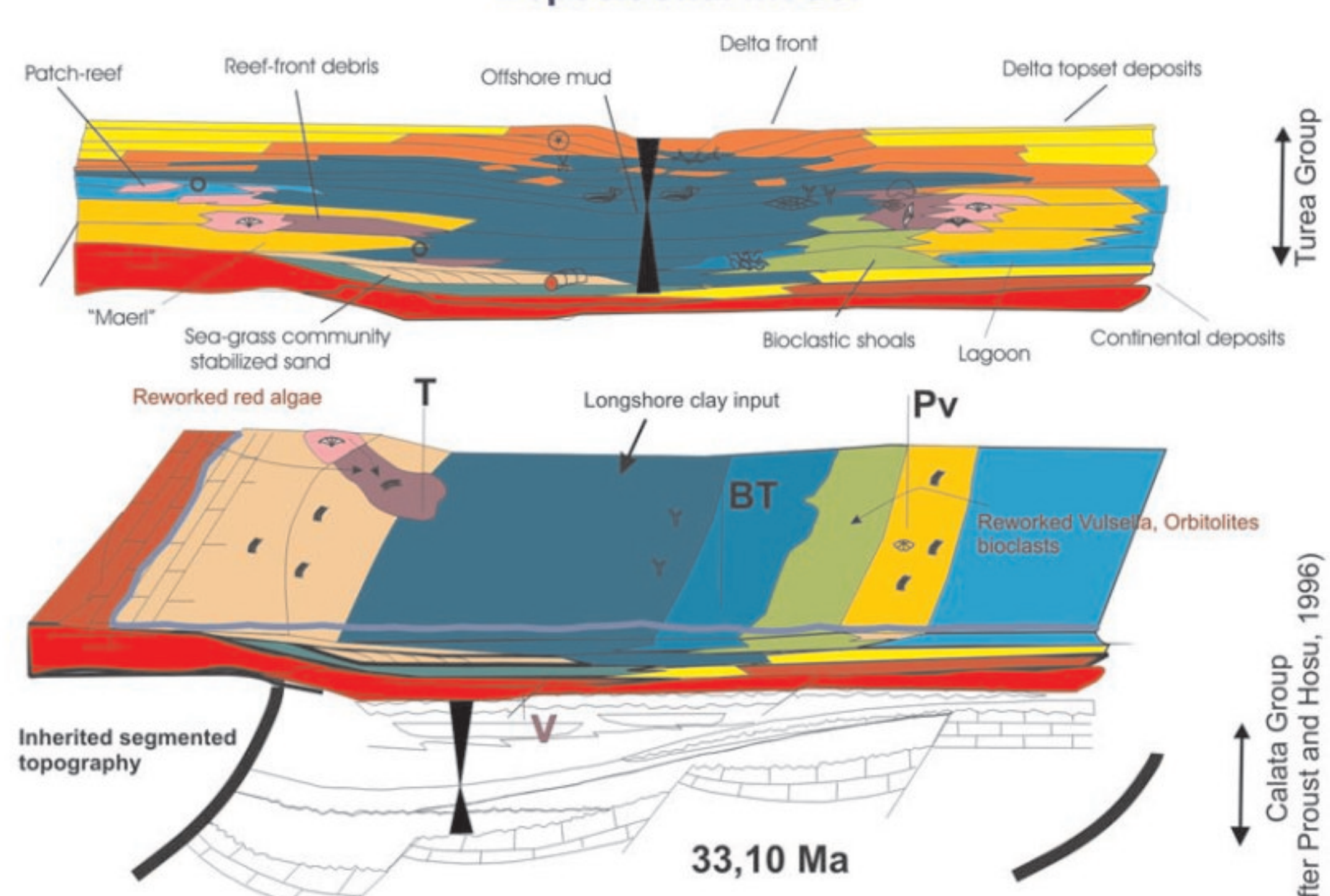
Cross-section

Synthetic curves of relative sea-level changes through the Turea Group depositional period (Priabonian-Earliest Rupelian) in the Transylvanian Basin



Localization of the studied profiles (A) and their stratigraphic range (B). Map and sedimentological log from Györfi et al. (1999) and Proust & Hossu (1996)

Depositional model



Methods

Each ecological assemblage and sedimentological paleoenvironment in shallow water shelf is directly or indirectly linked to water depth. The paleoenvironment is represented by a microfossils and sea-level changes correspond to changes in microfossils and qualitative/quantitative composition of bioclasts. Thus, a qualitative, quantitative and statistical analysis of bioclasts describes precisely paleoenvironments and is one of the most sensitive indicators of any paleoenvironmental changes, usually linked to relative sea-level changes or changes in accommodation space. Vertical fluctuations of each ecological assemblage percentage along a section reflect sea-level changes at that point and allows acquisition of relative paleobathymetry.

By this reasoning a quantitative paleoecological tool was outlined and applied successfully in different geological time-intervals, ranging from the Cretaceous of France (Arnaud-Vanneau, 1980) and Spain (Bernaus, 2000), to Paleocene (González-Lara, 2000).

In this work, counted data were acquired from thin-sections of carbonates and siliciclastics (average size of utile surface: 4cm x 6cm) and embedded artificial microfossils (average size of utile surface: 2cm x 2cm).

By developing the counting technique of 'thin-section mapping', counted specimens could be easily reviewed, and the database was quickly refreshed when needed.

In a preliminary report (Kovács and Arnaud-Vanneau, 2004), we defined 6 Paleocological Assemblages and one group made of echinoderm fragments for the Late Eocene-earliest Oligocene of the Transylvanian Paleogene Basin (Pv 2 section). These assemblages are composed of organisms constituting a biota, which lived more or less in the same paleoecological environments and were deposited at similar depth during that same time. After the analysis of additional outcrops (Sections BT, T, M, Pv1, Szp, V), these preliminary assemblages were further refined.

The quantification of relative seaward or landward shifting of the paleoecological assemblages allowed us to estimate the possible minimum and maximum paleodepth of deposition of the communities studied.

Interpretation of Sr^{87}/Sr^{86} isotopic data

33.10 True ages

37.10 Bioclasts reworked from the hinterland

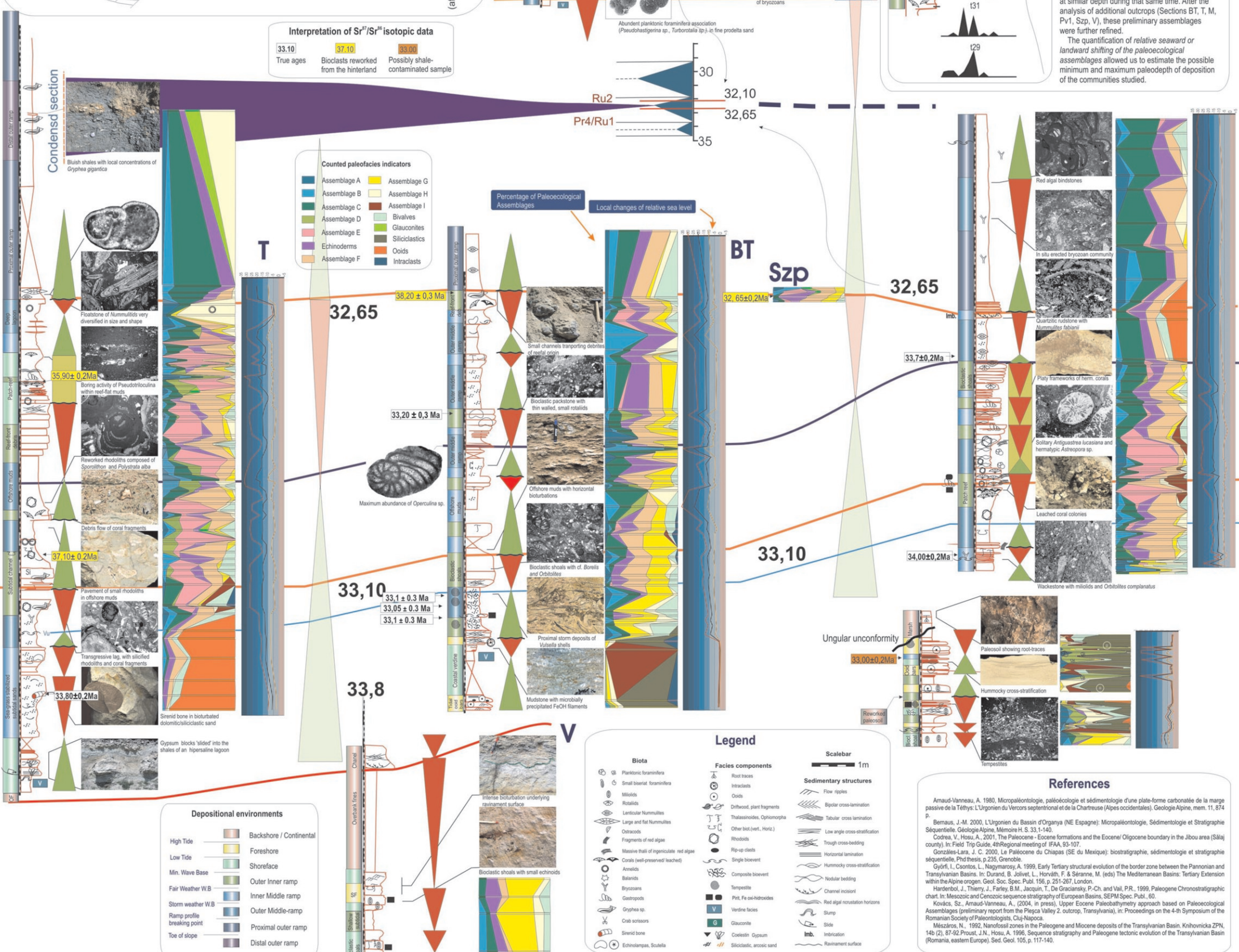
33.00 Possibly shale-contaminated sample

Counted paleofacies indicators

- Assemblage A
- Assemblage B
- Assemblage C
- Assemblage D
- Assemblage E
- Assemblage F
- Assemblage G
- Assemblage H
- Assemblage I
- Bivalves
- Glauconites
- Siliciclastics
- Ooids
- Intracasts

Percentage of Paleocological Assemblages

Local changes of relative sea level



Depositional environments

- High Tide
- Low Tide
- Min. Wave Base
- Fair Weather WB
- Storm weather WB
- Ramp profile breaking point
- Toe of slope
- Backshore / Continental
- Foreshore
- Shoreface
- Outer Inner ramp
- Inner Middle ramp
- Outer Middle-ramp
- Proximal outer ramp
- Distal outer ramp

Legend

- Biota: Planktonic foraminifera, Small bivalve foraminifera, Molluscs, Rotalids, Lenticular Nummulites, Large and flat Nummulites, Ostracods, Fragments of red algae, Massive thalli of irregularly calcified corals (well-preserved 'leached'), Corals (well-preserved 'leached'), Bryozoans, Gastropods, Gryphaea sp., Crab scapulae, Sponges, Echinodermata, Scapulae.
- Facies components: Root traces, Intertracasts, Ooids, Driftwood, plant fragments, Thalassiodonts, Ophiomorpha, Other biot. (vert. Horiz.), Rhizoids, Rip-up clasts, Single bioevent, Composite bioevent, Tempestites, Perv. Fe ox-hydroxides, Verdine facies, Glauconites, Coarsely Grained, Siliciclastic, arenaceous sand.
- Sedimentary structures: Flow ripples, Bipolar cross-lamination, Tabular cross lamination, Low angle cross-stratification, Trough cross-bedding, Hummocky cross-stratification, Nodular bedding, Channel incision, Red algal incision horizons, Slump, Side, Imbrication, Ravinement surface.

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