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GLOBAL STAGE	STAGE SLICE BASE	GENERALIZED $\delta^{13}\text{C}_{\text{org}}$ CURVE
HIRVANIAN	H2 end of HICE	
	K1a A. variabilis Zone (a)	
	K1b B. concolor Zone (a)	
KATIAN	K2 F. fauvel Zone (a)	
	K3 D. chonetes Zone (a)	
	K4 D. chonetes Zone (a)	
	K5 D. chonetes Zone (a)	
	K6 D. chonetes Zone (a)	
	K7 D. chonetes Zone (a)	
	K8 D. chonetes Zone (a)	
	K9 D. chonetes Zone (a)	
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	K12 D. chonetes Zone (a)	
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	K14 D. chonetes Zone (a)	
	K15 D. chonetes Zone (a)	
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	K18 D. chonetes Zone (a)	
	K19 D. chonetes Zone (a)	
	K20 D. chonetes Zone (a)	
	K21 D. chonetes Zone (a)	
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	K99 D. chonetes Zone (a)	
	K100 D. chonetes Zone (a)	
	K101 D. chonetes Zone (a)	
	K102 D. chonetes Zone (a)	
	K103 D. chonetes Zone (a)	
	K104 D. chonetes Zone (a)	
	K105 D. chonetes Zone (a)	
	K106 D. chonetes	

The new chronostratigraphic classification of the Ordovician System recently proposed by Bergström et al. (2009) has highlighted again the necessity of integrating different aspects of lithostratigraphy, biostratigraphy, chemostratigraphy and chronostratigraphy in order to identify with precision subdivisions of stratigraphical intervals. Regarding the Late Ordovician in particular, it is becoming apparent that identification of the $\delta^{13}\text{C}$ excursion (HICE) for recognition and subdivision of the Hirnantian Stage is essential. The stable isotopic values of carbon at the Cellon Section straddle around $+1\%$ throughout the Uggwa Limestone (increased from a value of -1.1 for carbonate in the underlying Uggwa Shale) and show a prominent excursion of $+2.8\%$ precisely at the unconformity with the overlying Plöcken Fm. (see Figure). If confirmed by high-resolution sampling this excursion coincides with the prominent peak in carbonate $\delta^{13}\text{C}$ at the Katian-Hirnantian boundary (HICE). Geochemical signals reveal a dynamic ocean chemistry during the Hirnantian in the Cellon Section (figure to the right). By using the ratio of highly reactive iron over total iron contents in the sediment, we get an estimation of the reducing conditions in the water column. The late Katian and earliest Hirnantian has unequivocal values, around the conventional threshold for anoxic values at 0.38, however above the Paleozoic mean for sediments deposited under oxic conditions. Moving into the Hirnantian and the Plöcken Fm., there is a clear enrichment of reactive iron. In the end-Hirnantian an increasing part of the reactive iron consists of pyrite. It seems a reducing water column, with increasing concentrations of sul

KATIAN

HIRNANTIAN

O/S boundary

Fe_{HR}/Fe_{TOT}

δ³⁴S_{Py}

δ¹³C (‰)

UGOWIA LIMESTONE FORMATION

PLOCKEN FORMATION

50 cm

1 **a**

2 **b**

3 **c**

4 **d**

5 **e**

6 **f**

7 **g**

8 **h**

9 **i**

10 **j**

11 **k**

12 **l**

13 **m**

14 **n**

15 **o**

16 **p**

17 **q**

18 **r**

19 **s**

20 **t**

21 **u**

22 **v**

23 **w**

24 **x**

25 **y**

26 **z**

27 **aa**

28 **ab**

29 **ac**

30 **ad**

31 **ae**

32 **af**

33 **ag**

34 **ah**

35 **ai**

36 **aj**

37 **ak**

38 **al**

39 **am**

40 **an**

41 **ao**

42 **ap**

43 **aq**

44 **ar**

45 **as**

46 **at**

47 **au**

48 **av**

49 **aw**

50 **ax**

51 **ay**

52 **az**

53 **ba**

54 **bb**

55 **bc**

56 **bd**

57 **be**

58 **bf**

59 **bg**

60 **bh**

61 **bi**

62 **bj**

63 **bk**

64 **bl**

65 **bm**

66 **bn**

67 **bo**

68 **bp**

69 **bq**

70 **br**

71 **bs**

72 **bt**

73 **bu**

74 **bv**

75 **bw**

76 **bx**

77 **by**

78 **bz**

79 **ca**

80 **cb**

81 **cc**

82 **cd**

83 **ce**

84 **cf**

85 **cg**

86 **ch**

87 **ci**

88 **cj**

89 **ck**

90 **cl**

91 **cm**

92 **cn**

93 **co**

94 **cp**

95 **cq**

96 **cr**

97 **cs**

98 **ct**

99 **cu**

100 **cv**

101 **cw**

102 **cx**

103 **cy**

104 **cz**

105 **da**

106 **db**

107 **dc**

108 **dd**

109 **de**

110 **df**

111 **dg**

112 **dh**

113 **di**

114 **dj**

115 **dk**

116 **dl**

117 **dm**

118 **dn**

119 **do**

120 **dp**

121 **dq**

122 **dr**

123 **ds**

124 **dt**

125 **du**

126 **dv**

127 **dw**

128 **dx**

129 **dy**

130 **dz**

131 **ea**

132 **eb**

133 **ec**

134 **ed**

135 **ee**

136 **ef**

137 **eg**

138 **eh**

139 **ei**

140 **ej**

141 **ek**

142 **el**

143 **em**

144 **en**

145 **eo**

146 **ep**

147 **eq**

148 **er**

149 **es**

150 **et**

151 **eu**

152 **ev**

153 **ew**

154 **ex**

155 **ey**

156 **ez**

157 **fa**

158 **fb**

159 **fc**

160 **fd**

161 **fe**

162 **ff**

163 **fg**

164 **fh**

165 **fi**

166 **fj**

167 **fk**

168 **fl**

169 **fm**

170 **fn**

171 **fo**

172 **fp**

173 **fq**

174 **fr**

175 **fs**

176 **ft**

177 **fu**

178 **fv**

179 **fw**

180 **fx**

181 **fy**

182 **fz**

183 **ga**

184 **gb**

185 **gc**

186 **gd**

187 **ge**

188 **gf**

189 **gg**

190 **gh**

191 **gi**

192 **gj**

193 **gk**

194 **gl**

195 **gm**

196

LATE ORDOVICIAN INTERVAL OF THE CELLON SECTION – Lithostratigraphic column based on new field measurements by HPS and AF. Vertical trends in key geochemical parameters (iron, sulphur and carbon isotopes) across the Hirnantian glacial event are shown; letters a-y indicate sampling points. Letters KKK indicate position of K-bentonite levels sampled for radiometric dating. New and revised biostratigraphic data indicate the standard *Normalograptus persculptus* Graptolite Zone, the *Amorphograptus ordovicicus* Conodont Zone, the *Tanuchina* *elmata* Chitinozoan Zone and the diagnostic *Himantia* brachiopod fauna. Trilobite faunas are also indicated (after Schönlaub et al., 2011).

New collections of graptolites, conodonts and chitinozoans have identified the index fossils for the global standard biostratigraphic zonations, are complementary to the faunal record documented previously and add a further recalibration of the latter biostratigraphic data. To date, the index graptolite for the lower Hirnantian, *Normalograptus extraordinarius* has not been found in the Carnic Alps. We conclude, however, that the siltstones of Member 2 of the Uggwä Limestone Fm. at the Cellon Section may correspond to this level (see Figure above) or that the index graptolite zone for the basal Hirnantian is encompassed in the unconformity there. Periglacial deposits which clearly reflect the diamictite nature of part of the Plöcken Fm. provide unequivocal evidence of the Hirnantian glaciation in this region. Further geochemical analyses and radiometric dating are in progress.

PERSULPTUS BIOZONE - Hirnantian graptolites of the Pilsbry Formation, and *Normalograptus persculptus* (Elles and Wood, 1907); a, b - Feilstritzgraben; a - transversally stretched specimen, b - longitudinally stretched specimen, c, d - Cellon Section; c - above bed 5, d - 85 cm above bed 4A, e - *Normalograptus ex gr. normalis* (Lapworth, 1877). Oberbuchach. Thick scale bar = 1 mm.

ACUMINATUS BIOZONE - Lowermost Silurian graptolites from black limestones and siliceous shales of Bischofalm Facies, a, b *Normalograptus normalis* (Lapworth, 1877), c, d, e *Normalograptus mirnyensis* (Obut and Sobolevskaya, 1967), e-f *Parakidograptus acuminatus* (Nicholson, 1867); f - transversally stretched specimen, f, i - longitudinally stretched specimens, j, k - *Glyptograptus aff. tamariscus* (Nicholson, 1968); l - *Neodidograptus bifurcus* (Ye, 1978), p-r *Neodidograptus laetus* Storch and Feist, 2008. Wasserfall Section, all specimens from bed 1 except for f from bed 2. Thick scale bar = 1 mm.

GRAPTOLITES

Fig. 1. *Tanuchitina elongata* (Gouché) from bed 7 in the Late Ordovician interval of the Celson Section. Total length of the vesicle: 373 μ m. Fig. 2. *Armoricocyclina nigerica* (Gouché) from bed 7 in the Late Ordovician interval of the Celson Section. Total length of the vesicle: 159 μ m. Fig. 3. *Desmochitina minor* Eisenack from bed 8 in the Late Ordovician interval of the Celson Section. Total length of the vesicle: 90 μ m. Fig. 4. *Tanuchitina elongata* (Gouché) from bed 5 in the Late Ordovician interval of the Celson Section. Total length of the vesicle: 246 μ m. Fig. 5. *Desmochitina minor* Eisenack from the sandy basalles above bed 8 in the Late Ordovician interval of the Celson Section. Total length of the vesicle: 79 μ m. Fig. 6. *Armoricocyclina nigerica* (Gouché) from the sandy basalles above bed 5 in the Late Ordovician interval of the Celson Section. Total length of the vesicle: 159 μ m.

BRACHIOPODS

CONODONTS

CHITINOZOANS

“there has been an urgent need for a globally applicable chronostratigraphic classification of the Ordovician” (Bergstrom et al., 2009).

